

The Crystal Structure of Leucophanite*

BY E. CANNILLO, G. GIUSEPPETTI AND V. TAZZOLI

Centro di Cristallografia del C.N.R., sez. VI, Istituto di Mineralogia dell'Università di Pavia, Italy

(Received 14 November 1966)

Leucophanite is a silicate of general formula $\text{CaNaBeSi}_2\text{O}_6\text{F}$. The space group assumed for this work is $P2_12_12_1$ with $a=7.401$, $b=7.420$, $c=9.939$ Å, $Z=4$. The true space group is probably $P1$. The crystal structure has been solved, on the basis of its analogy with that of another silicate, melilite, from a three-dimensional Patterson synthesis. The final bond distances and angles are obtained from coordinates refined by the least-squares method applied to three-dimensional orthorhombic data (final R value: 9.4%). Triclinic refinement is in progress. The crystal structure consists of infinite sheets of BeO_4 and SiO_4 tetrahedra between which Ca and Na atoms are arranged. The relationships between the crystal structures of melilite and leucophanite are described and a hypothesis is advanced to explain the differences between them.

Introduction

Leucophanite is a late-magmatic silicate containing Ca, Na, Be and F. Its chemical and crystallographic data, cited by Strunz (1966) in the last edition of *Mineralogische Tabellen*, are derived from two papers by Zachariasen (1930, 1931) on leucophanite and meliphanite (another silicate closely related to leucophanite). According to Zachariasen these minerals could both be derived from the 'ideal' compound $\text{Ca}_2\text{BeSi}_2\text{O}_7$. The former would be obtained by replacing $\frac{1}{2}\text{Ca}$ with Na and $\frac{1}{2}\text{O}$ with F, the latter by replacing $\frac{1}{3}\text{Ca}$ with Na, $\frac{1}{3}\text{Si}$ with Al and an appropriate amount of O with F.

Zachariasen also considered that meliphanite, being only a little different from the ideal compound, would probably have a crystal structure very near to that of melilite $\text{Ca}_2\text{MgSi}_2\text{O}_7$ (Warren, 1930), because of the similar properties of Be and Mg. In fact both minerals are tetragonal and have almost identical cell parameters (taking into account the different orientation and the doubling of c in meliphanite). On the other hand, the replacement of O by F in leucophanite would exceed the threshold at which a distortion in the structure would occur producing a lower symmetry, in this case orthorhombic.

The structural model proposed by Warren was successively confirmed and refined by Smith (1953). The crystal structures of both meliphanite (Dal Negro, Rossi & Ungaretti, 1967) and leucophanite have been analysed in our laboratory.

Experimental

For the present work natural crystals from Brevig (Norway) were used. Two chemical analyses carried out by us are in substantial agreement with the published formula: $\text{CaNaBeSi}_2\text{O}_6\text{F}$.

The lattice parameters were determined at room temperature by measuring with a single-crystal diffractometer the 2θ values of appropriate reflexions and extrapolating the corresponding interplanar spacings, plotted against $\frac{1}{2}(\cos^2\theta/\sin\theta + \cos^2\theta/\theta)$, to $2\theta=180^\circ$. The values so obtained are: $a=7.401 \pm 0.008$, $b=7.420 \pm 0.008$, $c=9.939 \pm 0.005$ Å.

The theoretical specific gravity computed on the basis of these cell parameters for $Z=4$ is 2.961, whereas the experimental one is 2.959 g.cm^{-3} .

The space group used for this work is $P2_12_12_1$ according to the following systematic absences: $h00$ with $h=2n+1$, $0k0$ with $k=2n+1$, $00l$ with $l=2n+1$ and to Zachariasen's considerations. Actually not only three $h00$ reflexions with $h=2n+1$ were detected, but a careful inspection of the X-ray pictures revealed slight differences in the intensities of hkl , $\bar{h}kl$, $hk\bar{l}$ and $h\bar{k}l$ reflexions. These differences are significantly larger than the errors in measuring the intensities. So the true space group would be $P1$. However, it has been thought that the differences mentioned above are due to such small displacements from the orthorhombic symmetry as to be not essential in the determination of the crystal structure. For this reason the intensities of only one octant of the reciprocal sphere were measured. Of course the question must be reconsidered in the refinement.

A cylindrical specimen of leucophanite was used for recording intensities. The axis of the cylinder was parallel to the b axis and the radius was 0.016 cm. Weissenberg equi-inclination photographs of hkl (k from 0 to 6) reflexions were taken with nickel-filtered copper radiation, using the multiple film technique. A total of 628 independent reflexions in the space group $P2_12_12_1$ was inspected; 42 of them did not give any observable blackening on the films.

The intensities were measured with a Nonius microdensitometer and corrected for α_1 - α_2 spot doubling and for Lorentz-polarization and absorption factors. The linear absorption coefficient is $\mu=147$ cm^{-1} for Cu $K\alpha$ radiation and the absorption factors $A^*=A^{-1}$ ranged

* Paper presented at the VII Congress of the International Union of Crystallography, Moscow, USSR, July 1966.

from 7.5 to 51.5. A secondary extinction correction was applied during refinement.

The intensities of the different levels were approximately on the same scale owing to the equal exposure times. After each structure factor calculation the scale was improved by the criterion $\Sigma F_o = \Sigma F_c$, applied separately to each level.

Structure analysis

Patterson sections normal to [010] were computed and the determination of the crystal structure was carried out on the basis of the following considerations: (1) A comparison of the intensities of melilite with those of leucophanite revealed some relevant correspondences. (2) The cell parameters of melilite are: $a=7.789$, $c=5.018$ Å; if these values are compared with those of leucophanite, one can see that a and b of this mineral are nearly identical with a of melilite while c has twice the length. (3) An arrangement of cations in positions analogous to those occupied by cations in melilite therefore appeared very probable, since these positions have a symmetry consistent also with the space group $P2_12_12_1$. Furthermore such positions should be occupied, at least to a certain degree, by different cations in the two halves into which the leucophanite cell may be divided (c of leucophanite is twice c of melilite) in order to account for its lower symmetry. (4) An analysis of the relative heights of the maxima of convenient Patterson sections should permit the determination of the differences in the arrangement of the atoms in the two halves of the cell and consequently the kinds of atoms occupying each position.

Some different arrangements derived from the Patterson sections were examined by means of their effects on calculated electron density projections on a plane normal to [010]. The correctness of the chosen hypothesis was verified by means of a three-dimensional electron density synthesis in which all the atoms were detected. At this stage the discrepancy factor computed on all the observed reflexions was 0.36.

Refinement

The method of least squares was applied in order to minimize the unweighted differences $F_o - F_c$. The calcu-

lations were carried out with a program which takes into account the full matrix (Sgarlata, 1965); the structure factors were derived from the atomic scattering factors obtained from Moore's (1963) constants considering all the atoms as half-ionized. At first three cycles were accomplished assuming an isotropic tem-

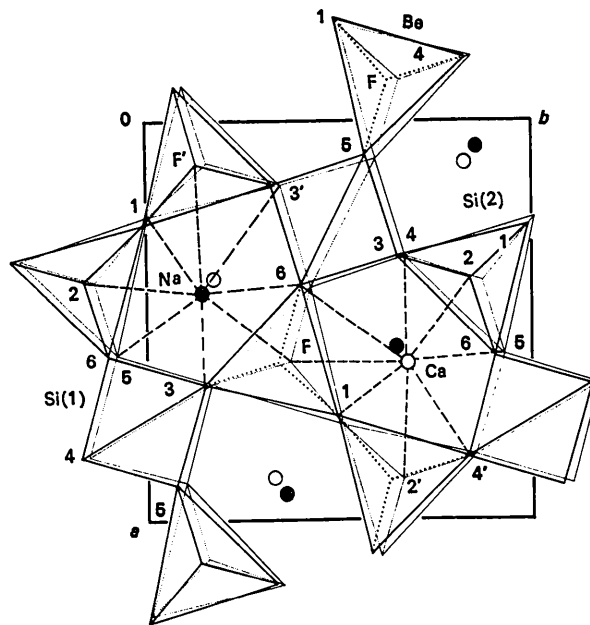


Fig. 1. Projection of the unit cell and key to the identification of the atoms and their coordination. The numbers without other specification refer to oxygen atoms.

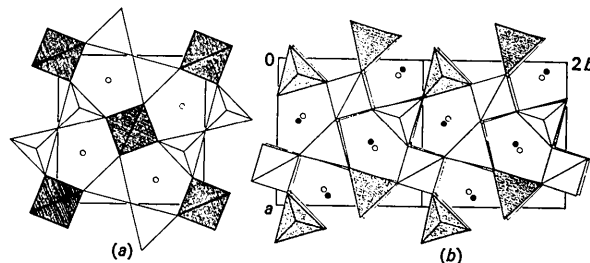


Fig. 2. (a) The crystal structure of melilite projected along [001]. (b) The crystal structure of leucophanite. Open circles: Ca; filled circles: Na; white tetrahedra: Si; hatched tetrahedra: Mg; dotted tetrahedra: Be.

Table 1. Atomic coordinates and their estimated standard deviations (fractional $\times 10^4$); isotropic temperature factors

	x/a	y/b	z/c	B
Ca	0.1050 (4)	0.8265 (5)	0.9894 (3)	1.23 Å ²
Na	0.0670 (11)	0.8535 (13)	0.5059 (8)	2.33
Si(1)	0.7406 (5)	0.0132 (7)	0.2507 (4)	0.85
Si(2)	0.1080 (5)	0.1508 (7)	0.2265 (4)	0.71
Be	0.1091 (24)	0.1268 (50)	0.7174 (32)	1.84
O(1)	0.2421 (20)	0.0037 (23)	0.1623 (14)	1.38
O(2)	0.1052 (20)	0.1607 (29)	0.3829 (22)	1.46
O(3)	0.8382 (12)	0.8484 (21)	0.3381 (12)	1.28
O(4)	0.8383 (14)	0.8308 (20)	0.8476 (12)	1.33
O(5)	0.9130 (14)	0.0746 (19)	0.6574 (10)	1.46
O(6)	0.9087 (17)	0.1000 (18)	0.1627 (13)	1.35
F	0.1021 (13)	0.1315 (17)	0.8753 (9)	1.66

Table 2. Structure factors of leucophanite ($\times 10$)

Reflexions marked with an asterisk were unobservably weak; in these cases F_0 is derived from $\frac{1}{2}I_0$ min.

Table with columns for h, k, l, F0, Fc, A, B and multiple rows of data. The table is organized into several sections based on hkl values, including 'h 0 1', 'h 1 1', 'h 3 1', 'h 4 1', and 'h 5 1'. Each row contains numerical values representing structure factors for different reflections. Some values are marked with an asterisk (*).

perature factor for each atom; the progress of refinement was indicated by a fall of the R value from 0.36 to 0.12.

Four more cycles with the attribution of anisotropic thermal factors to each atom followed by secondary extinction correction lowered R to 0.086. However this anisotropic refinement was not considered completely reliable because it was accomplished on data scaled by layers and only on the basis of the independent reflexions of the orthorhombic space group. Moreover the standard deviations of all the parameters are exceedingly high and thus seem to indicate that the parameters obtained for an atom (and its equivalents) in the space group $P2_12_12_1$ are an average of the independent values that could be obtained in the space group $P1$. However, these values would not be very different from those published here.

The anisotropic refinement will be carried out with all the independent triclinic reflexions. At the present time the refinement has been brought to completion as follows: a secondary extinction correction was calculated after the third isotropic least-squares cycle because a plot of $\ln I_c/I_o$ against I_c for the 104 most intense reflexions ($I_o > 1000$ on the absolute scale) indicated a conspicuous secondary extinction effect. A linear relation between $\ln I_c/I_o$ and I_c was assumed and a least-squares secondary extinction coefficient derived. This was applied as a correction to all the F_o 's through the formula: $(F_o)_{corr} = (F_o)_{ext} \cdot \exp(\frac{1}{2}\epsilon I_c)$ where $\epsilon =$

$3.739 \cdot 10^{-5}$. The above correction lowered the R value to 0.097. After two more isotropic cycles, when the shifts were largely less than the standard deviations, the refinement was stopped.

The final R for 586 observed reflexions was 0.094 and 0.100 for all the 628 reflexions (giving to the unobserved ones a value half of the least observed intensity).

The atomic parameters, their standard deviations (Cruickshank, 1949) and the isotropic thermal parameters are shown in Table 1. The observed and calculated structure factors are presented in Table 2. Interatomic distances and angles are presented in Table 3 as well as their standard deviations (Ahmed & Cruickshank, 1953; Darlow, 1960). Reference should be made to Fig. 1 to identify the labelled atoms.

Discussion

The crystal structure of leucophanite is closely related to that of melilite. The two crystal structures are compared in Fig. 2. In melilite Si_2O_7 groups are present and are connected by MgO_4 tetrahedra to form sheets between which the calcium atoms are arranged. The sites occupied by Ca in the melilite structure are split in leucophanite into two series of positions occupied by Ca and Na respectively. Beryllium does not occur in the Mg site but in one of the Si sites, and one Si occupies the Mg site while the other lies in the same position as in melilite. So no Si_2O_7 group is present in

Table 3. *Interatomic distances, angles, and their standard deviations (in brackets)*

	Bond length	Mean		Bond angle	Mean
Si(1)O(5)	1.605 (11) Å		O(5)Si(1)O(4)	112.3 (6)°	
O(4)	1.615 (14)		O(6)	112.7 (6)	
O(6)	1.651 (13)		O(3)	108.1 (7)	
O(3)	1.664 (14)	1.634 Å	O(4)Si(1)O(6)	108.0 (7)	
			O(3)	111.8 (7)	
Si(2)O(2)	1.556 (22)		O(6)Si(1)O(3)	103.6 (6)	109.4°
O(1)	1.607 (16)		O(2)Si(2)O(1)	115.9 (9)	
O(3)	1.649 (15)		O(3)	110.5 (9)	
O(6)	1.649 (13)	1.615	O(6)	112.5 (7)	
BeO(1)	1.565 (32)		O(1)Si(2)O(3)	107.5 (7)	
F	1.570 (33)		O(6)	104.1 (7)	
O(5)	1.616 (24)		O(3)Si(2)O(6)	105.6 (6)	109.3
O(4)	1.691 (38)	1.610	O(1)BeF	112.7 (1.9)	
CaO(2)	2.354 (19)		O(5)	110.8 (2.2)	
O(5)	2.374 (13)		O(4)	104.9 (1.6)	
O(1)	2.389 (15)		FBeO(5)	110.1 (1.5)	
O(2')	2.393 (16)		O(4)	111.7 (2.1)	
O(4)	2.425 (11)		O(5)BeO(4)	106.3 (2.0)	109.4
F	2.531 (12)				
O(4')	2.639 (12)				
O(6)	3.032 (13)				
NaO(1)	2.352 (17)				
O(3)	2.377 (13)				
F	2.381 (14)				
O(5)	2.501 (14)				
O(6)	2.525 (15)				
O(2)	2.601 (23)				
F'	2.773 (12)				
O(3')	2.945 (14)				

leucophanite, but zigzag chains of SiO_4 tetrahedra are formed and are nearly parallel to [010]. These chains are connected by BeO_4 tetrahedra to form sheets between which the atoms of Ca and Na lie. If, as Zoltai (1960) correctly suggests, no distinction is made between SiO_4 , BeO_4 and MgO_4 tetrahedra, leucophanite can be classified as a single sheet silicate together with melilite (and meliphanite). The above mentioned chains could justify the {100} and {101} cleavages while the {001} cleavage is due to the presence of the sheets.

The differences between the crystal structures of melilite and leucophanite can be due, according to Zachariassen, to the presence of fluorine in the latter. In fact in synthetic Be-akermanite (Goria, 1953) and gugiaite $\text{Ca}_2\text{BeSi}_2\text{O}_7$ (Peng, Tsao & Chou, 1962) there is a complete replacement of Mg by Be and presumably Be occupies the position of Mg and is connected with four Si atoms. The tetragonal symmetry is not modified: the space group of Be-akermanite and gugiaite is $P4_2/m$; the cell parameters of the former are $a=7.501$, $c=4.931$ Å and of the latter $a=7.48$, $c=5.04$ Å. But if in leucophanite Be were to remain in the positions of Mg in melilite, fluorine would be connected with Si, and this does not generally happen in natural silicates.

The bond distances and angles of the tetrahedral cations are quite regular and comparable to those found in other silicates.

The bond distances of Ca and Na are listed in Table 3 in order of increasing length. It is rather difficult to decide what is the coordination number of these cations on the basis of these bond distances. Smith describes the coordination polyhedron of Ca in melilite as an archimedean antiprism built up by four atoms at about 2.44 Å and four at about 2.70 Å, but the resulting electrostatic equilibrium is unsatisfactory however one computes it. In the present case one can reject the eight-coordination for Ca and, perhaps, the seven-coordination for Na, because some distances are surely

too long. Moreover a careful analysis shows that the interaction between Ca and Na and the seventh atom must be very little because in both cases this atom is nearly completely screened by other oxygen atoms. The sixfold coordination for both the cations gives also the least discrepancies in the electrostatic equilibrium. The sixfold coordination polyhedron of Ca is not an octahedron but resembles a trigonal prism; the length of its sides ranges from 2.65 to 3.65 Å. The sixfold coordination polyhedron of Na is a very distorted octahedron (sides ranging from 2.62 to 4.54 Å, O-Na-O angles ranging from 58 to 157°). For these reasons leucophanite does not seem to be referable to Belov's (1961) 'second chapter' silicates. However, no conclusive statement can be made before the triclinic refinement of the crystal structure.

References

- AHMED, F. R. & CRUICKSHANK, D. W. J. (1953). *Acta Cryst.* **6**, 385.
 BELOV, N. V. (1961). *Crystal Chemistry of Large-Cation Silicates* (English translation), p. 3. New York: Consultants Bureau, 1964.
 CRUICKSHANK, D. W. J. (1949). *Acta Cryst.* **2**, 65, 154.
 DAL NEGRO, A., ROSSI, G. & UNGARETTI, L. (1967). *Acta Cryst.* **23**, 260.
 DARLOW, S. F. (1960). *Acta Cryst.* **13**, 683.
 GORIA, C. (1953). *Atti Accad. Torino, Classe sci. fis. mat. nat.* **88**, 153.
 MOORE, F. H. (1963). *Acta Cryst.* **16**, 1169.
 PENG, C.-S., TSAO, Y.-L. & CHOU, T.-Y. (1962). *Scientia Sinica*, **11**, 977.
 SGARLATA, F. (1965). *Period. Min.* **34**, 401.
 SMITH, J. V. (1953). *Amer. Min.* **38**, 643.
 STRUNZ, U. (1966). *Mineralogische Tabellen*, p. 346. Leipzig: Akad. Verlag Ges. Geest & Portig K.-G.
 WARREN, B. E. (1930). *Z. Kristallogr.* **74**, 131.
 ZACHARIASEN, W. H. (1930). *Z. Kristallogr.* **74**, 226.
 ZACHARIASEN, W. H. (1931). *Norsk geol. Tidsskr.* **12**, 577.
 ZOLTAI, T. (1960). *Amer. Min.* **45**, 960.